

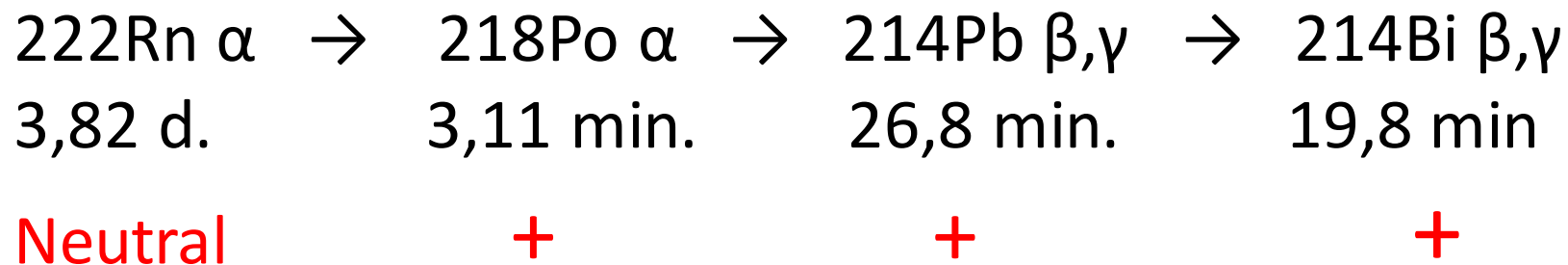
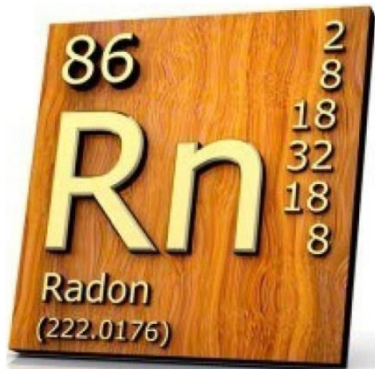
Estimation of volumetric activity of
Radon (^{222}Rn) at Aragats outdoors
and indoors (SKL) using the intensity
of the Bi214 line (609 keV)

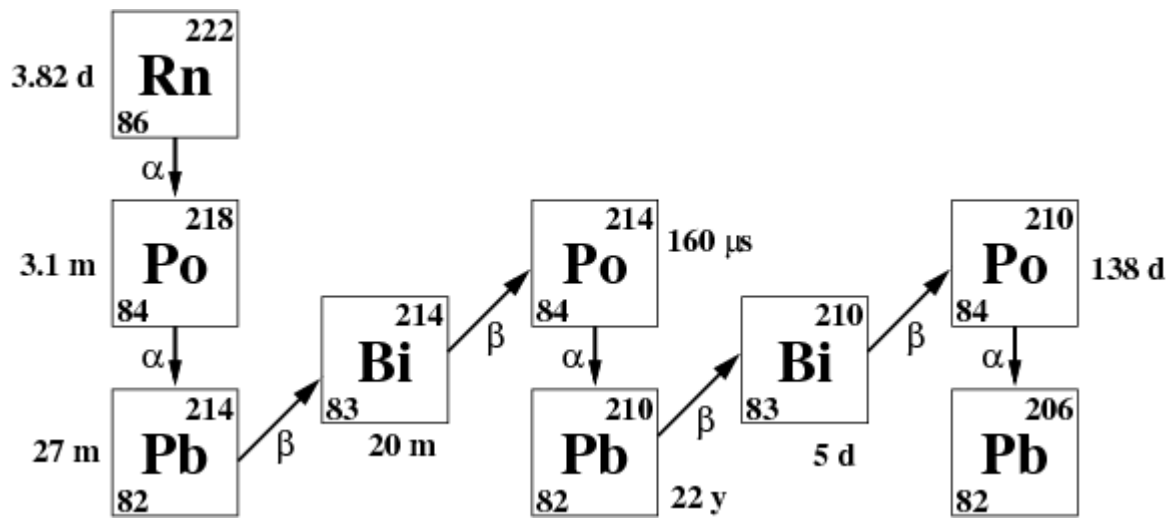
Eduard Mnatsakanyan

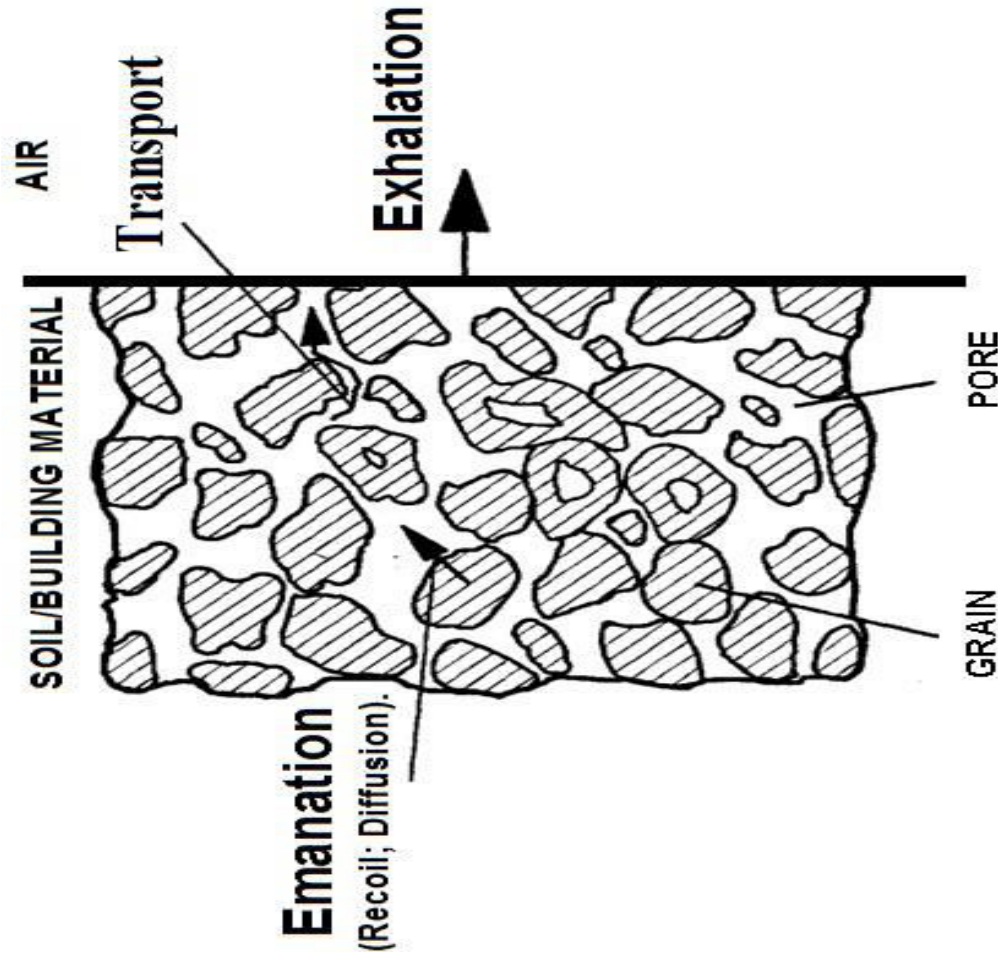
A.I. Alikhanyan National Science Laboratory (YerPhI)

1. All daughter nuclides of radon-222 are either positively charged or neutral and cannot significantly rise into the upper atmosphere under conditions of alternating electric fields.
2. On the other hand, exponential results show that growth of the gamma photons independent of the polarity of the alternating (perturbed) electric field.
3. as shown by mathematical calculations, the transfer of nuclides from the lower layers to the upper layers leads to a decrease in the count of gamma photons.

The excess of gamma photons can apparently be explained by a change in meteorological conditions







Exhalation Rate---[Bq m⁻².s⁻¹] ;

Volumetric activity in the
Atmosphere-----[Bq m⁻³].

The average **exhalation rate** of radon from the ground is **5–50 mBq/m²s**, leading to a near ground level radon concentration of 1–10 Bq/m³, but varies widely with ground conditions

Concentrations of radon in the outdoor environment are affected by exhalation rates from the ground and atmospheric mixing phenomena. In **turbulent conditions -turbulent mixing**. Stable atmospheric **stratification (temperature inversions)**, such as typically occur in daytime, much more vertical mixing occurs than at night, when inversions can trap radon close to the ground. This means that outdoor radon concentrations can vary diurnally by more than a factor of ten. There are also seasonal variations, related to the effects of precipitation, or to changes in prevailing winds. **Typical long-term average outdoor radon concentrations are about 10 Bq/m³** each for radon-222 and radon-220. However, long-term average concentrations of radon-222 vary from approximately **1** to more than **100 Bq/m³**, with the former typical of isolated **small islands** and the latter typical of sites with high radon exhalation rates over **large areas**.

. Moreover, progeny concentrations are affected by the depletion due to fallout, washout, rainout, and other scavenging effects. It is known that the physical processes mentioned above are dependent on certain meteorological parameters specific to the planetary boundary layer: air temperature, soil temperature, atmospheric pressure, wind speed, precipitations and relative humidity.

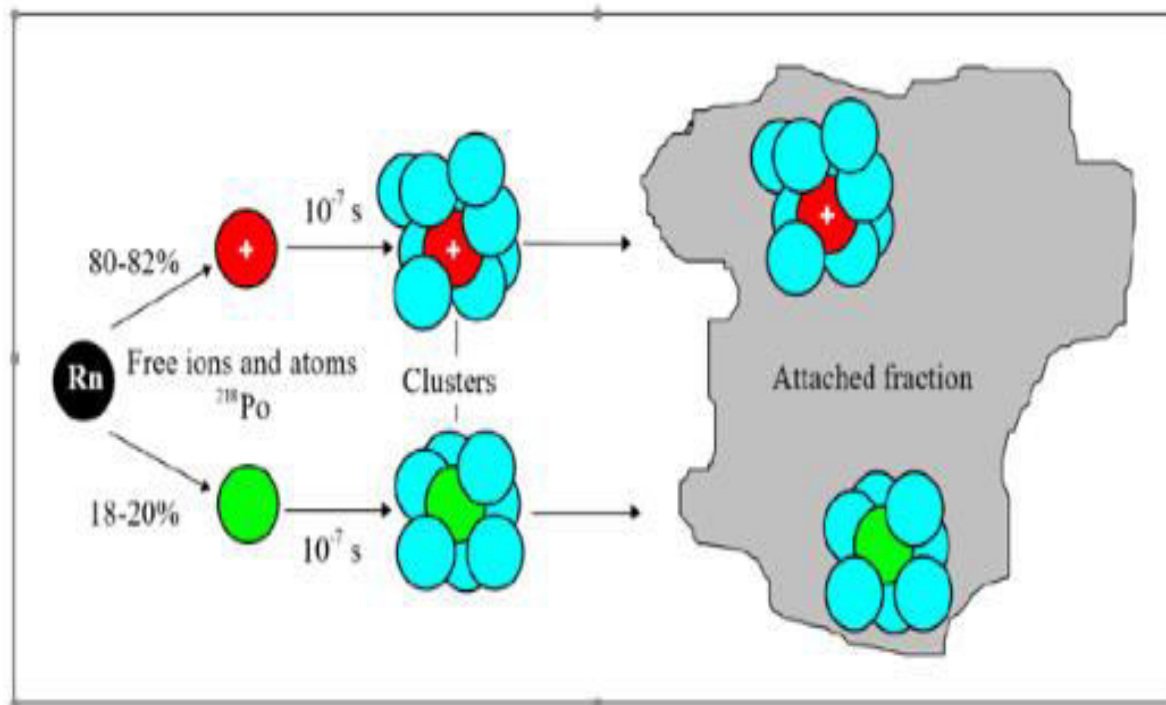


Fig. 1. Basic processes of radon decay product behaviour in air defining the concept of "unattached" and "attached" fractions [1]

[1] J.Porstendörfer, Properties and behavior of radon and thoron and their decay products in the air. J. Aerosol Sci. 25(1994)

Secular equilibrium

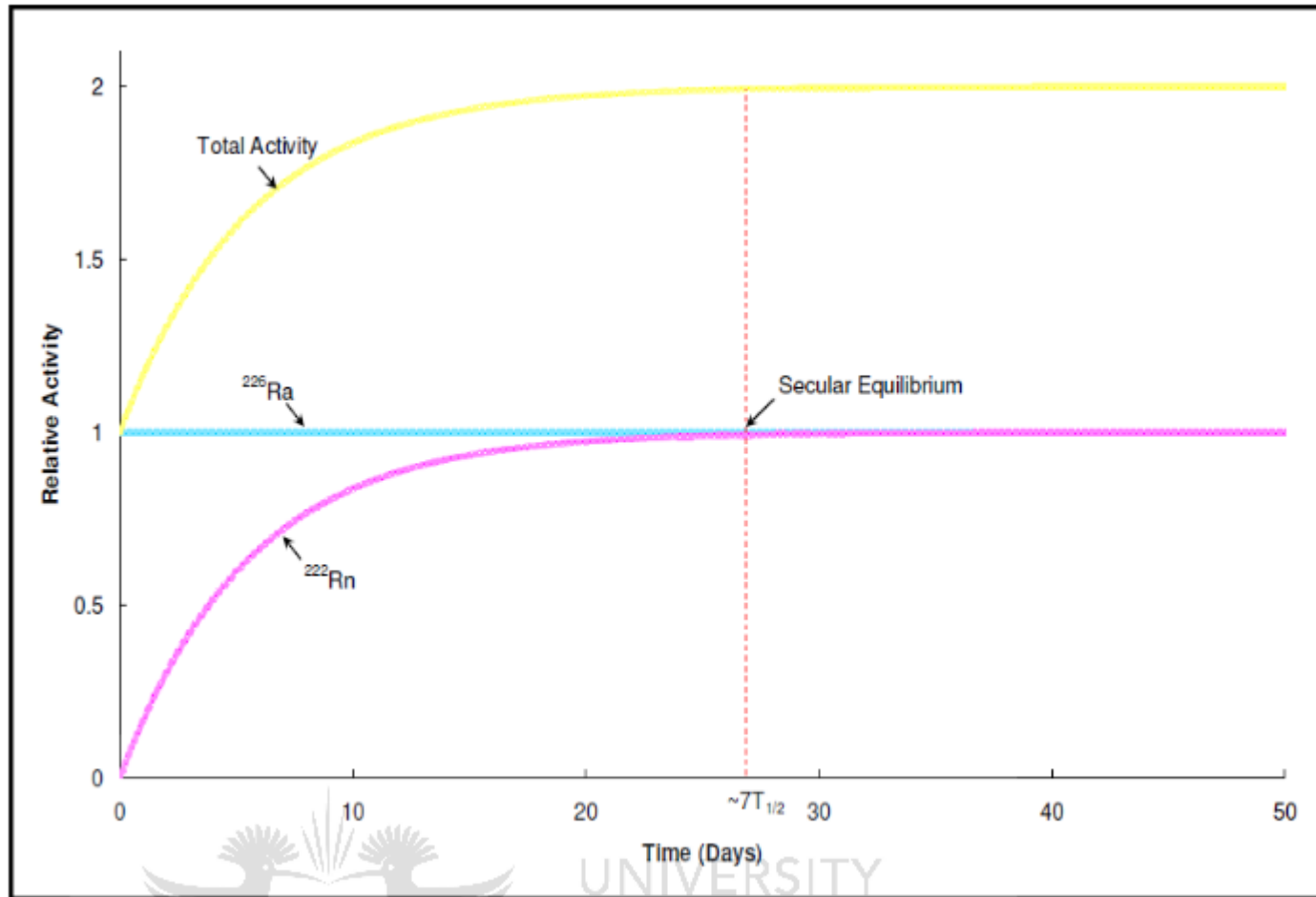


Figure 2.2 Secular equilibrium between ^{226}Ra and its daughter, ^{222}Rn

Transient equilibrium

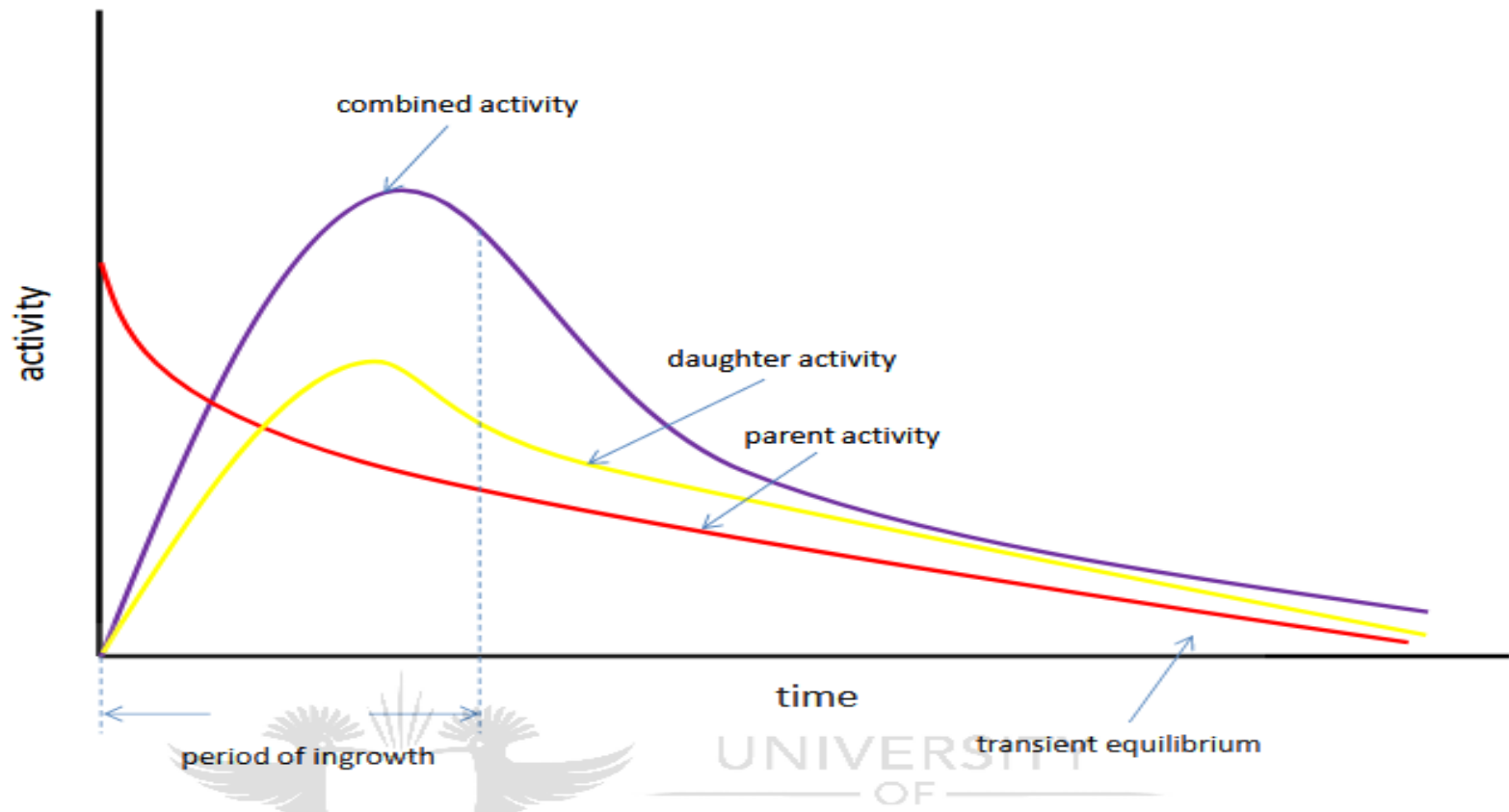


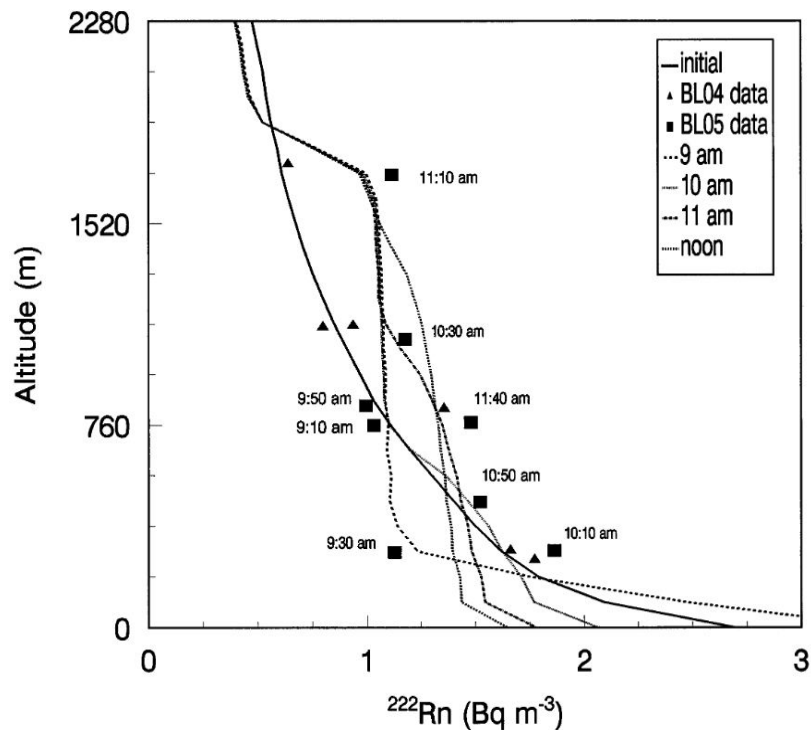
Figure 2.3 An illustration of transient equilibrium.

As time progresses, the parent, daughter and (parent + daughter) activities can be characterized by essentially the same decay constant (this is evident from Equation 2.16 for $t \rightarrow \infty$), as illustrated in Figure 2.3. [The Accurate...]

H, м	^{222}Rn , %	H, м	^{220}Th , %
0,01	100	0	100
1,00	95	5	70
10,0	87	10	50
100	69	25	20
1000	38	50	5

Табл. 6. Изменение концентрации радона и торона с высотой

^{222}Rn		^{220}Rn	
Высота, м	Содержание, %	Высота, м	Содержание, %
0,01	100	0	100
1	95	5	70
10	87	10	50
100	69	25	20
1000	38	50	5
	7	100	0,5



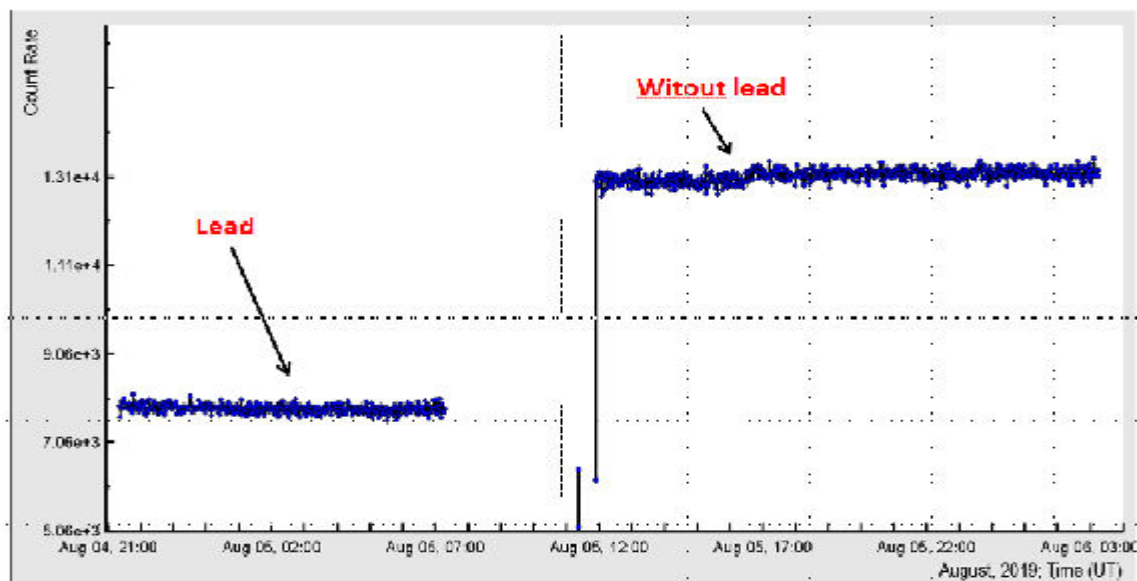
Based on the above data take approximately

$$A(h) = A_0 e^{-\lambda h};$$

where A_0 - Volumetric Activity of ^{222}Rn at surface layer;

$$\lambda = 0.0015 \text{ m}^{-1};$$

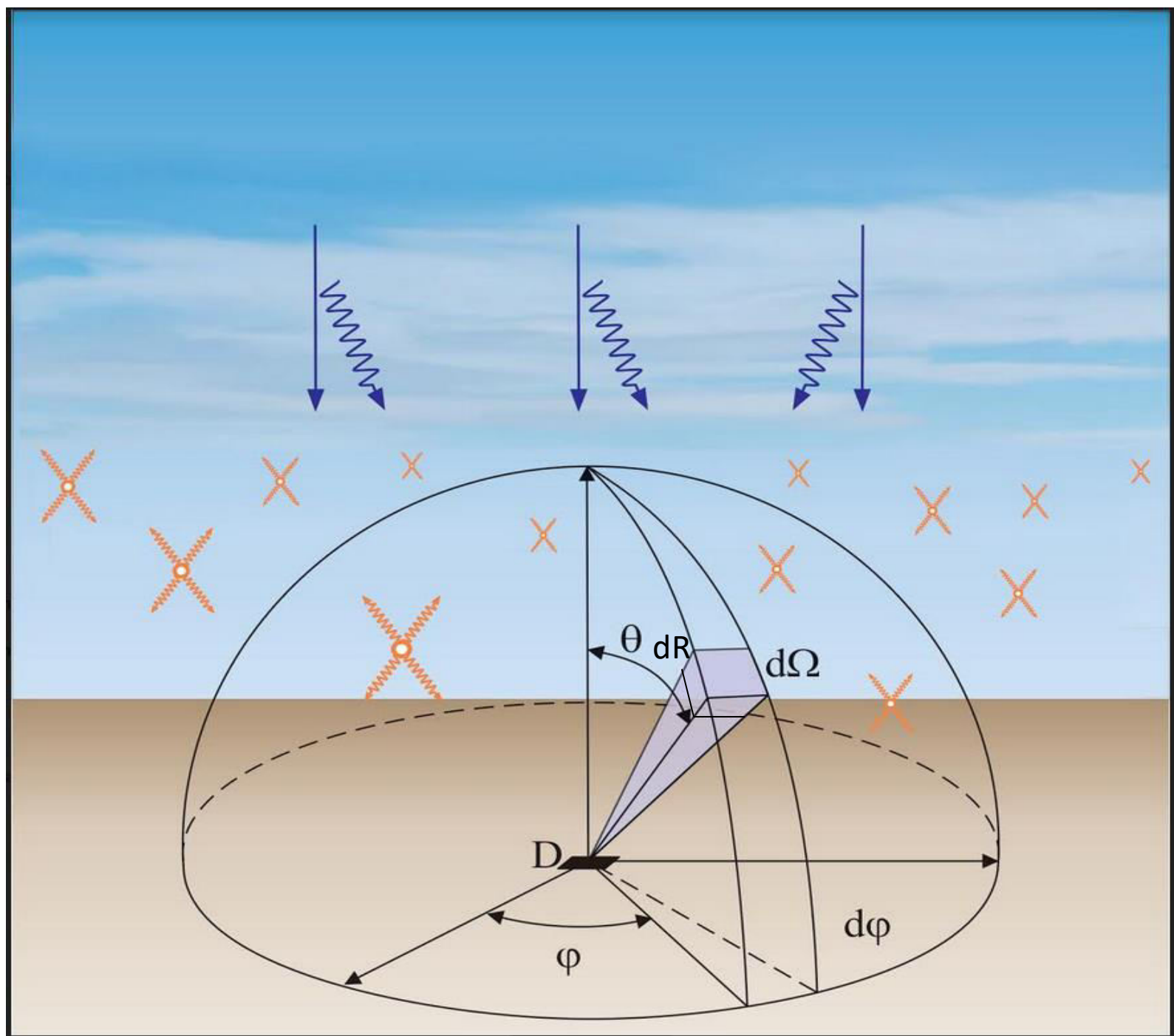
Det. ORTEC NaI(Tl)



Tab. 05 August 2019 det. ORTEC NaI(Tl)

1h

	Without lead	Lead	Difference	%
0.26 -10 MeV	779110	467282	311828	40
0.33 -0.38 MeV	23613	11898	11715	49,6
0.56 -0.66 MeV	26007	14193	11814	45,4
1 -1.2 MeV	5790	3216	2574	44,5
1.34 - 1.57 MeV	17387	9868	7519	43,2
1.62 - 1.9 MeV	2216	1046	1170	52,8



Direct measurement of the concentrations of all short-lived decay products of ^{222}Rn and ^{220}Rn are difficult and limited. They are estimated from considerations of equilibrium (or disequilibrium) between these nuclides and their respective decay products

By integrating in spherical coordinates and by taking into account the azimuthal symmetry of the model, the Count Rate (/) 609keV photons emitted by atmospheric ^{214}Bi and registered by the NaI(ORTEC) is given by the following equation:

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$$I = \frac{1}{4\pi} K A_0 S \int_{\varphi=0}^{2\pi} \int_{\theta=0}^{\pi/2} \int_{R=0}^{\infty} e^{-\lambda R \cos \theta} e^{-\mu R} \sin \theta dR d\theta d\varphi;$$

Where: I -is Count Rate [s^{-1}];

$$K = K_{eff} K_p;$$

K_{eff} -- intr. eff. ORTEC at 609keV (^{214}Bi)

K_p -- relative intensity of 609keV;

A_0 -- Volumetric Activity of ^{222}Rn [Bq/m^3], at ground level

S - Detector area [m^2];

R - Distance [m];

μ - Attenuation coefficient for dry air [m^{-1}];

λ - Height Attenuation coefficient [m^{-1}];

θ - Zenith Angle; φ - Azimuth angle;

For:

$$K=K_{eff} K_p = 0.38 \times 0.47;$$

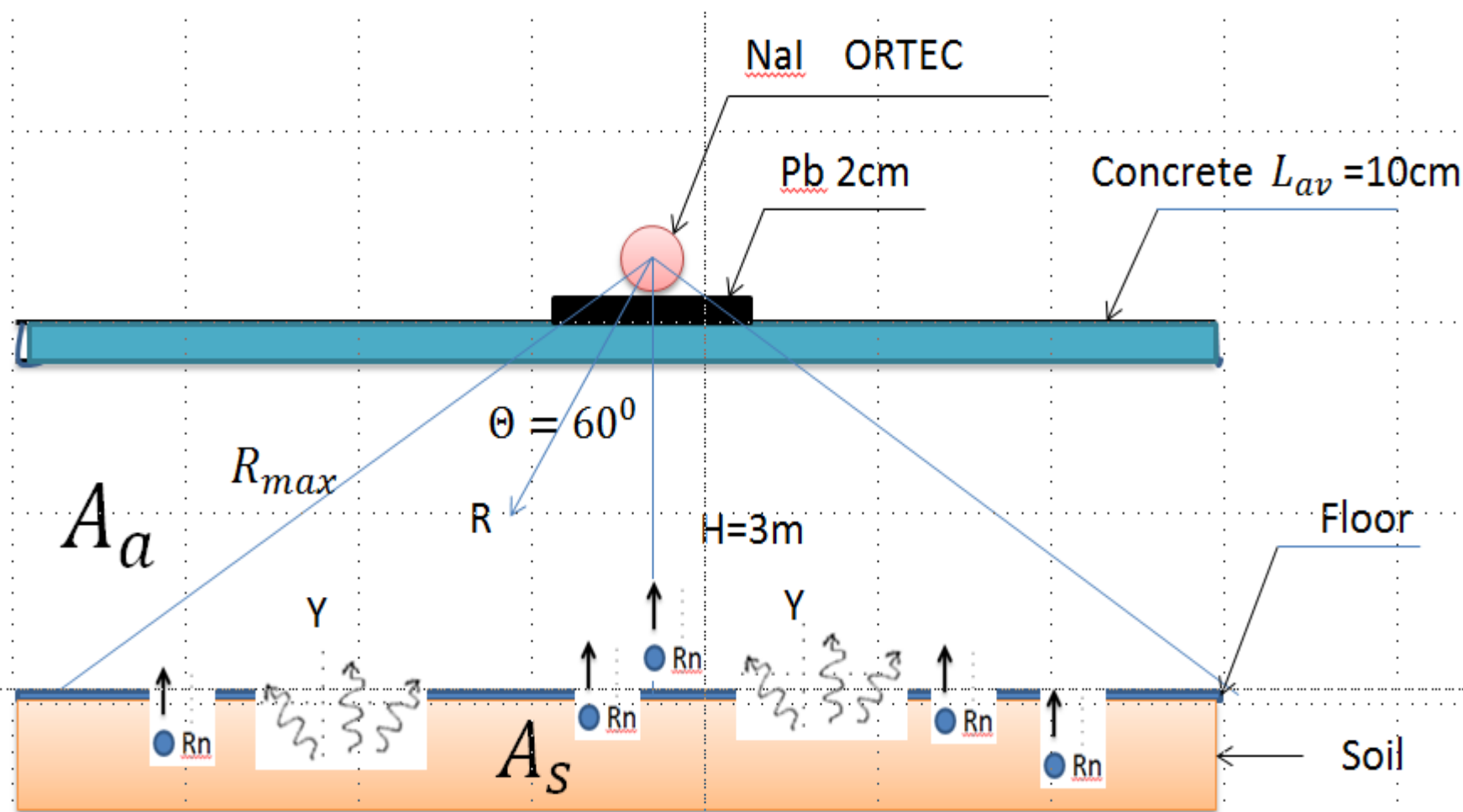
$$S(ORTEC)= 0.005 \text{ m}^2;$$

$$\mu=0.0696 \text{ m}^{-1};$$

$$\lambda =0.0015 \text{ m}^{-1};$$

$$l= 14193 h^{-1}= 3.94 c^{-1}. \text{ We get:}$$

$$A_0= 72 \text{ Bq/m}^3.$$



Typical radon emanation rates from the earth's surface:

- $2 \cdot 10^{-4}$ Bq / m²s for rocks,
 - (8-21) $\cdot 10^{-3}$ Bq / m²s for mountain soils,
 - (4-50) $\cdot 10^{-3}$ Bq / m²s for ??soils,
 - (5-38) $\cdot 10^{-3}$ Bq / m²s for desert soils (21-53) $\cdot 10^{-3}$ Bq / -m²s for black soils.
 - In areas with anomalous geology and geochemistry of underlying rocks, the rate of radon exhalation reaches $5.25 \cdot 10^2$ Bq / m²s.
- Weighted average over the area of the earth exhalation rate of radon $1.6 \cdot 10^{-2}$ Bq / m²s

The content of radon in soil air varies from $2.6 \cdot 10^3$ to $4.4 \cdot 10^4$ Bq / m³

The Russia measurements -In general, the values of ^{222}Rn in the surface layer of the atmosphere, over the territory of Russia, vary from 0.5 to 75 Bq / m³

- Rn flux from the soil --- 0.014 to 0.097 Bq / m²s with maximum values in the mountainous regions of the Urals, Southern Siberia and the Far East. average value is 0.040 Bq / m²s. The measurements were taken from a mobile laboratory which was part of a passenger train moving along the Trans-Siberian Railway from Moscow to Vladivostok.

Based on the above data for our calculations, select- $A_s = 44$ kBq/m².

For:

$$K=K_{eff}K_p = 0.38 \times 0.47;$$

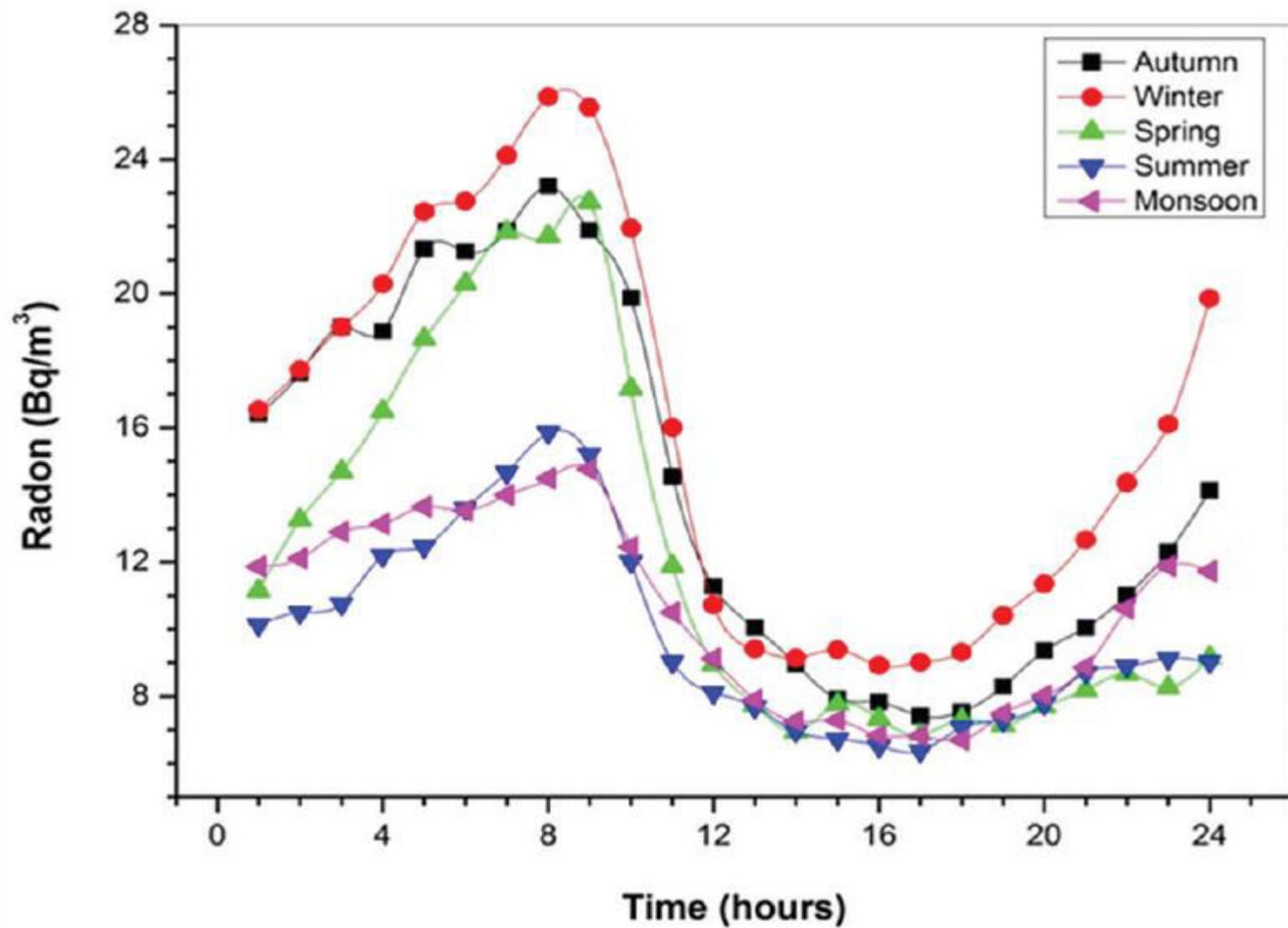
$$S_{(ORTEC)} = 0.005 \text{ m}^2;$$

$$\mu(\text{ordinary concrete}) = 0.08236 \text{ cm}^2/\text{g}; \text{---} 0.16472 \text{ cm}^{-1};$$

$$A_a = 72 \text{ Bq/m}^3;$$

$$A_s = 44 \text{ kBq/m}^3. \text{ We get:}$$

$$1) - I = 20 \text{ h}^{-1}; \text{ 2) } - I = \underline{9442 \text{ h}^{-1}}. (\text{By ORTEC} = \underline{11814 \text{ h}^{-1}}).$$



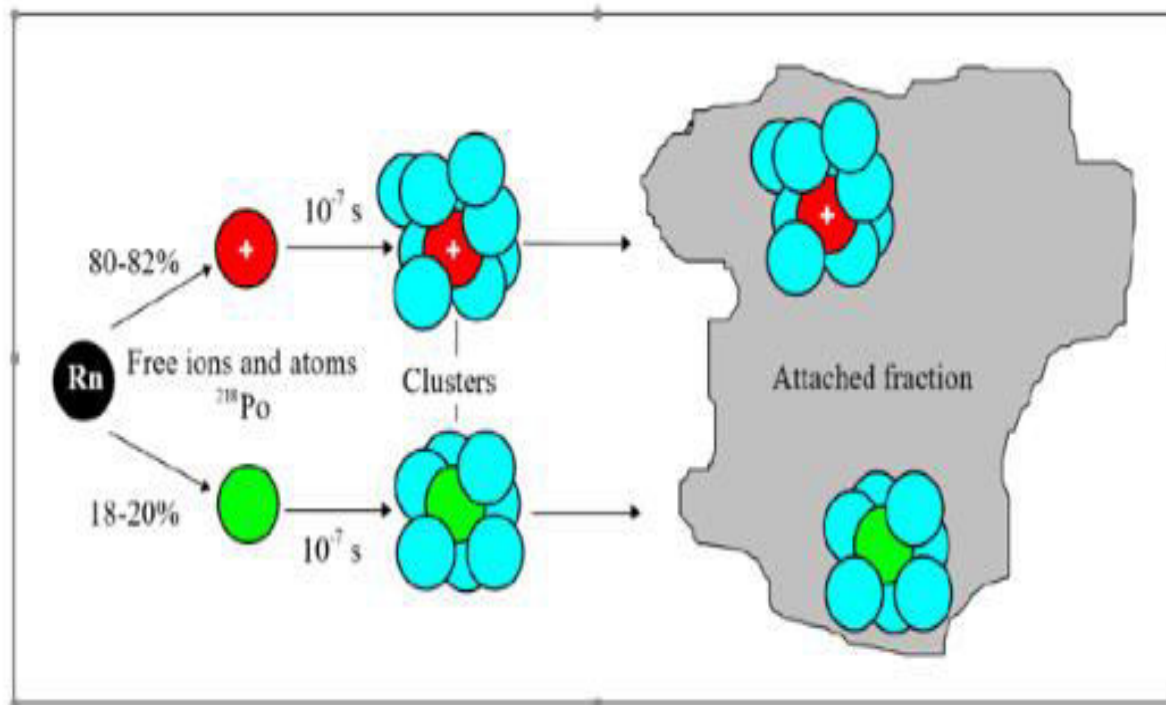
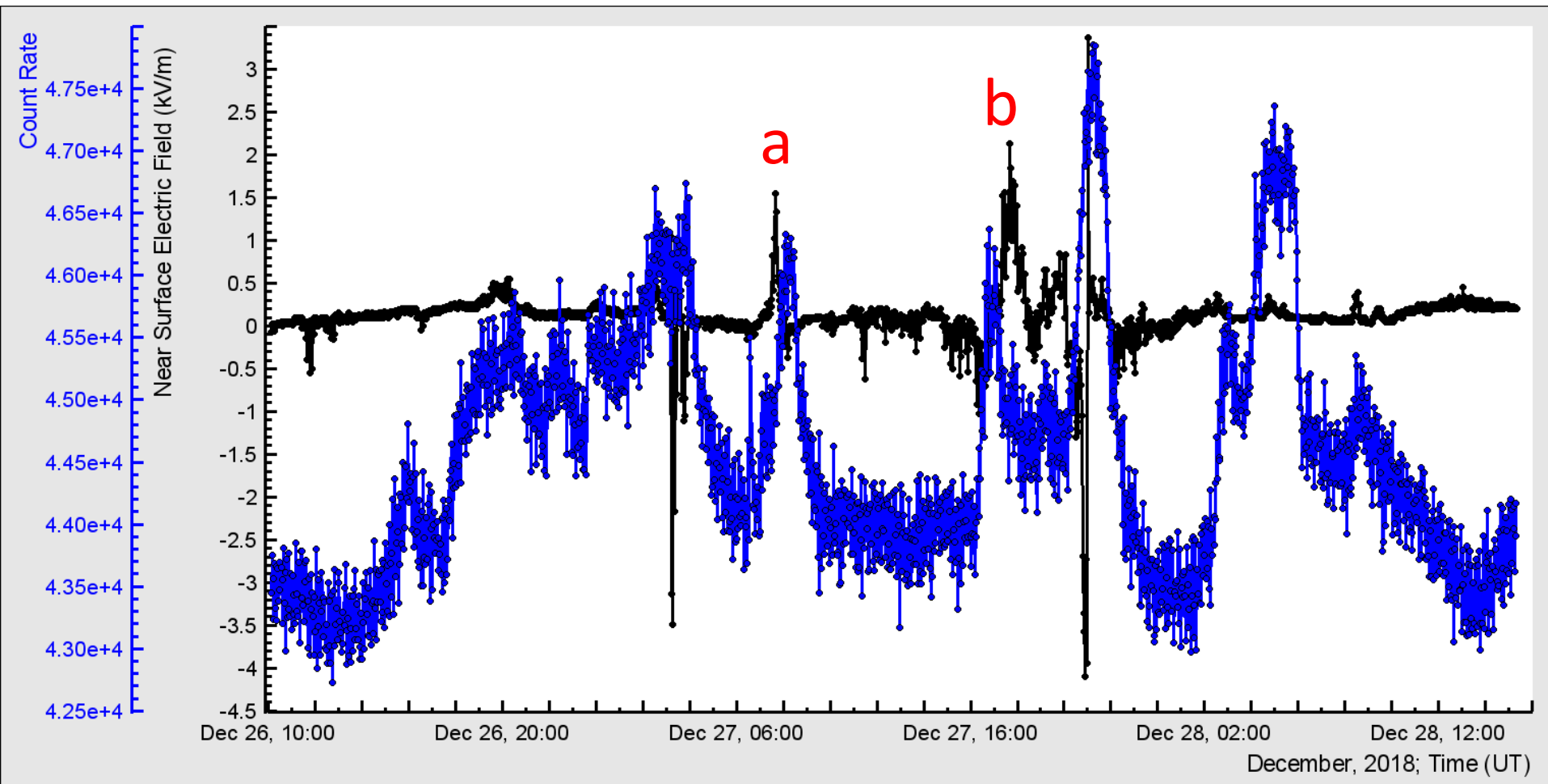


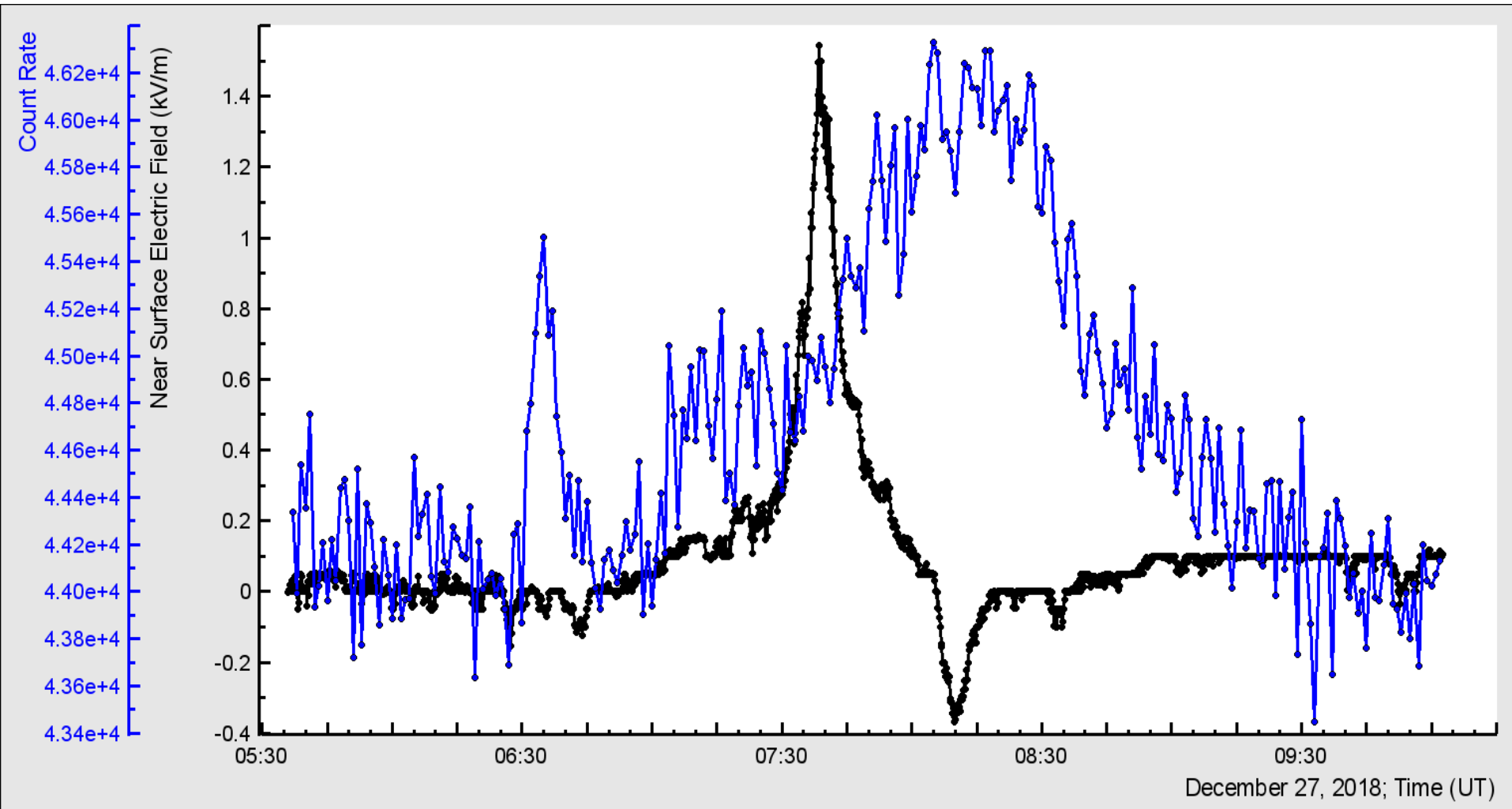
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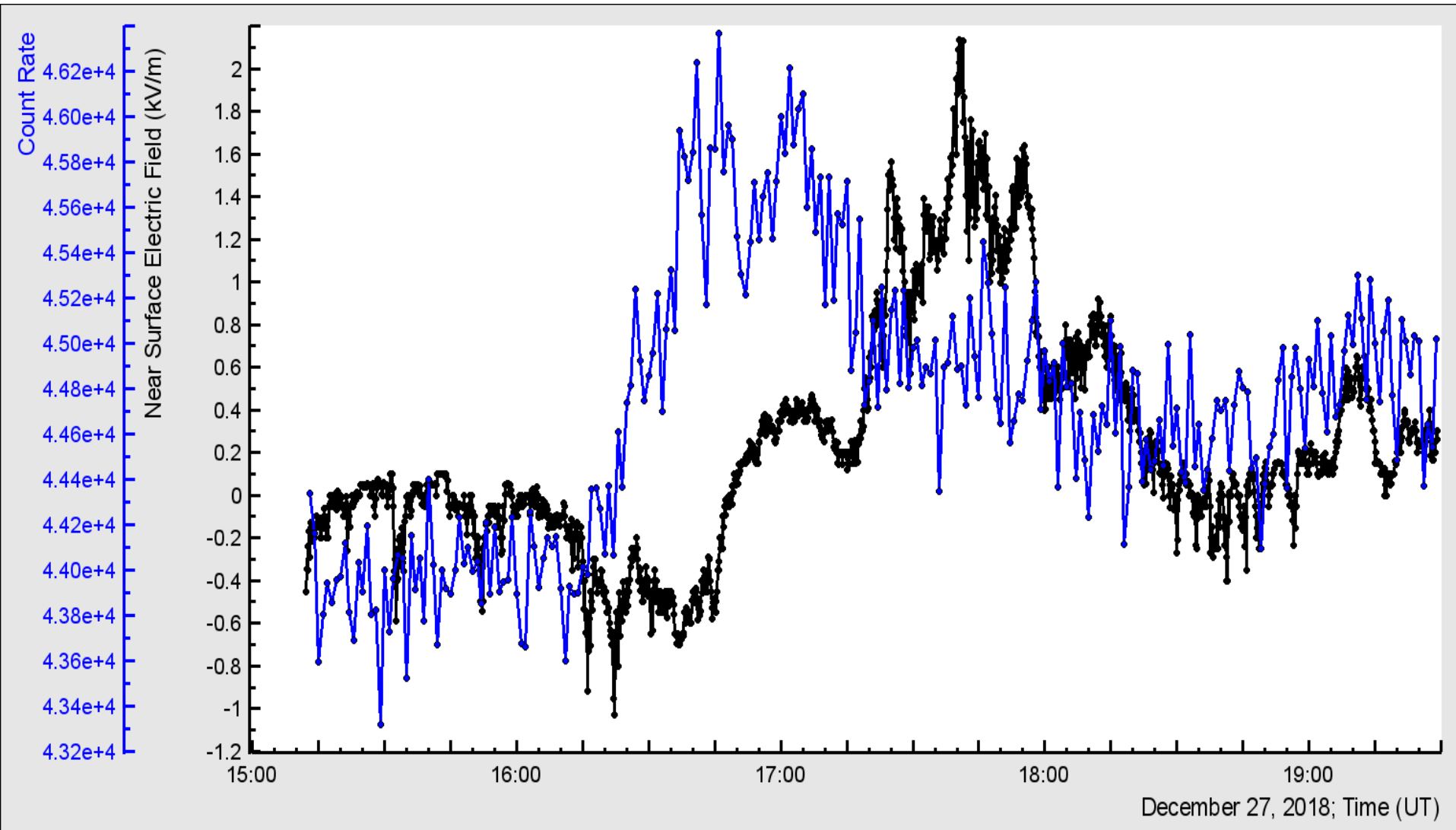
Dec.



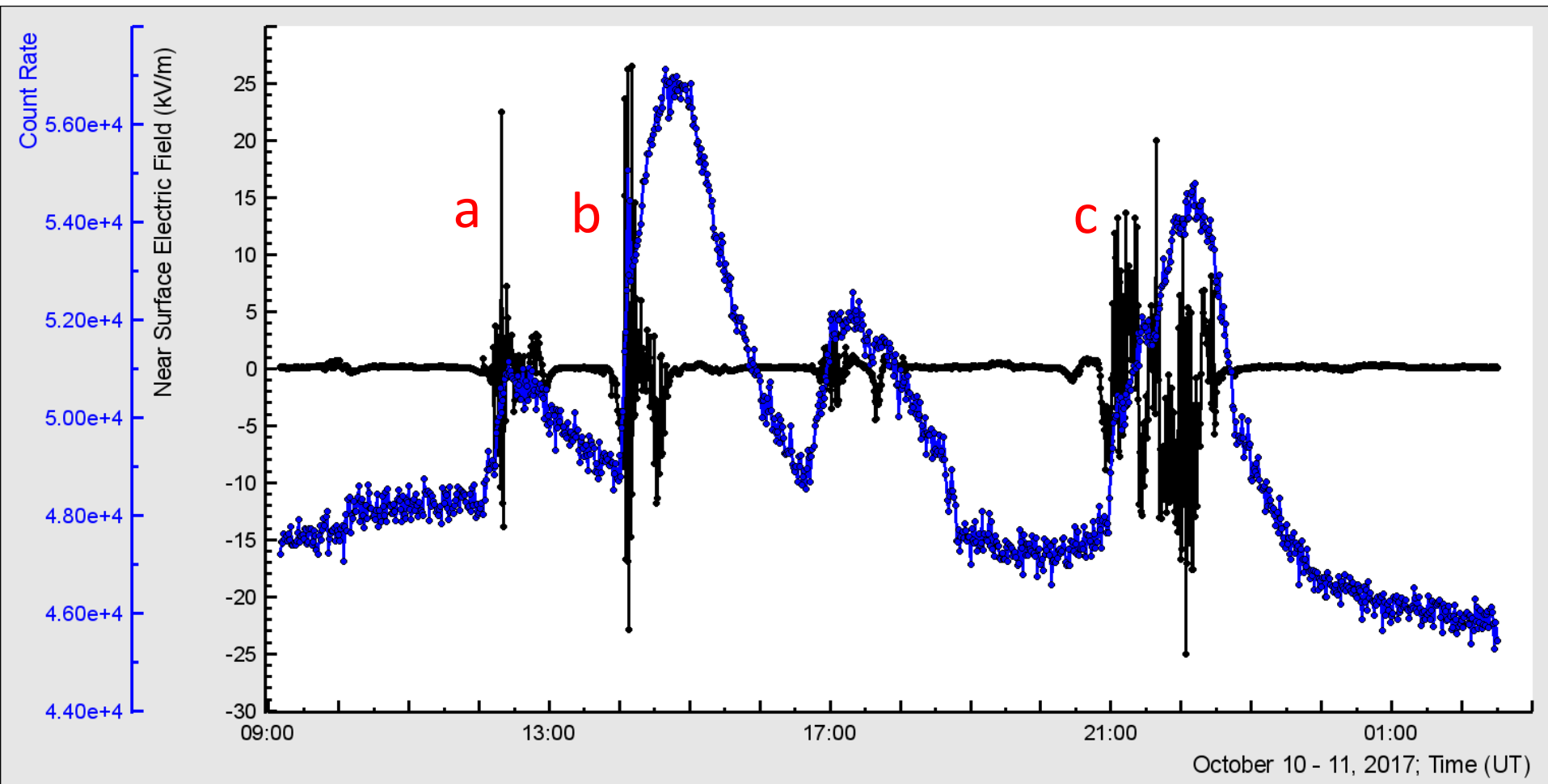
Dec. a



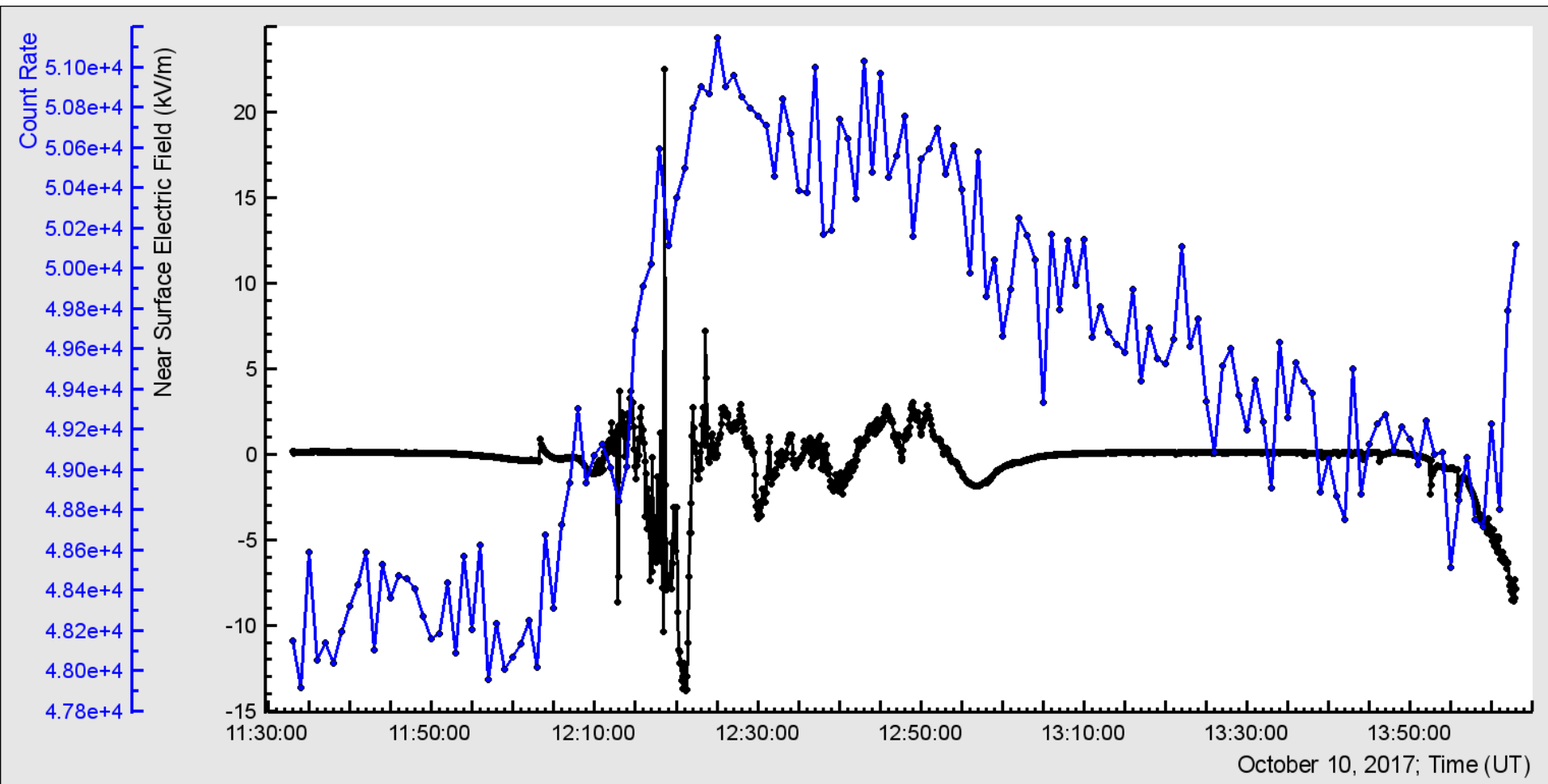
Dec. b



Oct.

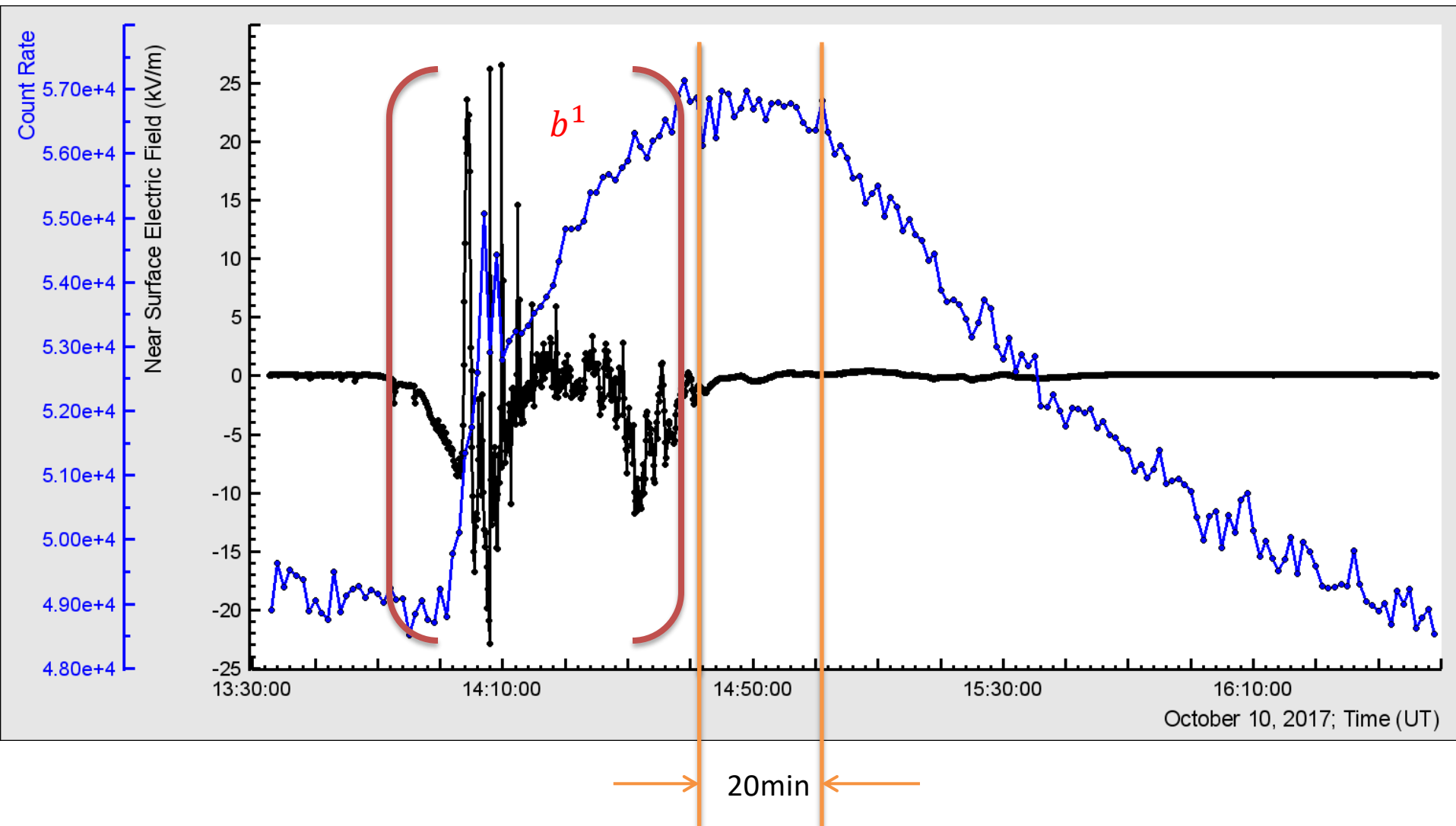


Oct. a

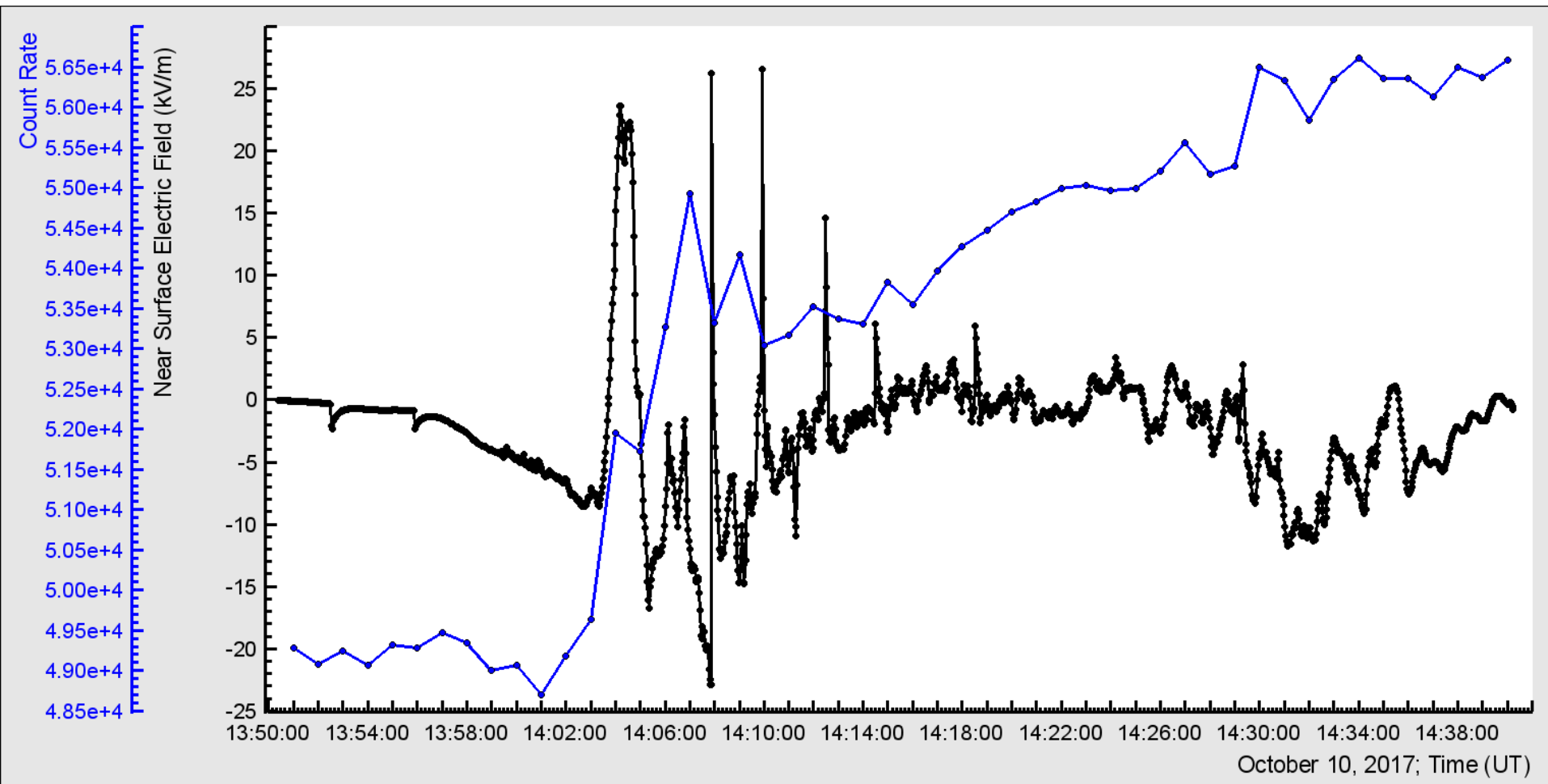


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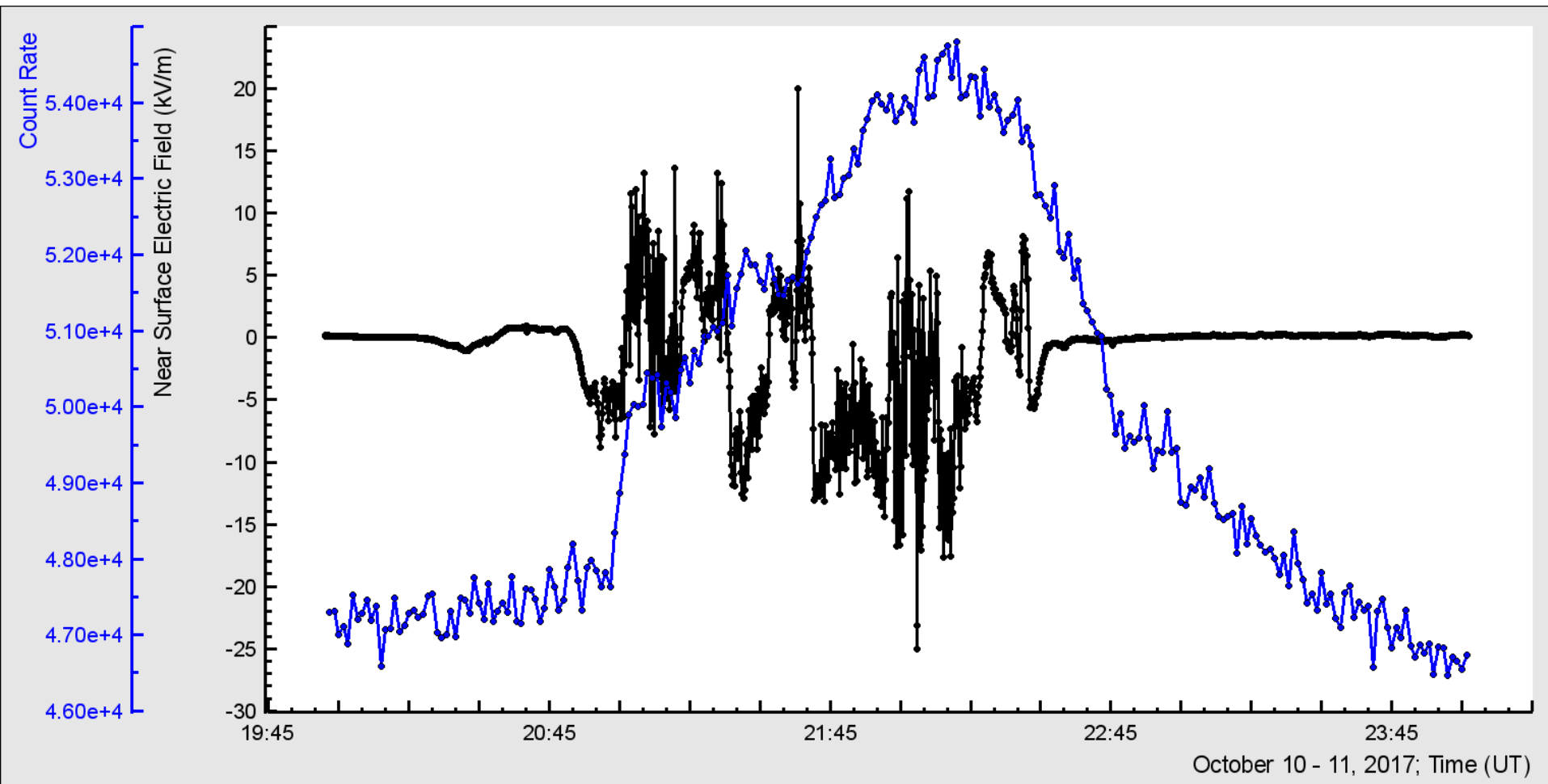
b



b^1



Oct. C



Conclusion

1. Most of the gamma coming from closer horizontal directions.!
2. A significant part (40%) of the gamma falling on the detectors (under the roof SKL) comes directly from the underlying soil.
3. Any changes gamma photon flux (if A_{ex} is constant) possible with vertical “compression” of the Radon and its daughter nuclides vertical distribution
4. Or A_{ex} changing by weather conditions.

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Thank you for attention